

GEOMORPHOLOGIC ANALYSIS OF FORILLON NATIONAL PARK, QUEBEC, CANADA

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INTRODUCTION

This paper summarizes the geomorphology of Forillon National Park, Quebec. The work is primarily a photo geologic interpretation in conjunction with a literature review conducted during the period March 20 to April 20, 1978.

A geomorphic map at a scale of 1:100 000 accompanies the paper and for the first time presents a comprehensive inventory of principal geomorphic features of the park classified according to age and process. Also, the mapping has developed new information with regard to glacial and pre-glacial geomorphic features. Despite a great many past studies in the park few have concentrated on the geomorphology and, therefore, little quantitative data exists.

It is important to point out that this work has been completed without the benefit of a field check. A considerable amount of field work would be necessary to map the park adequately and to establish the ground truth on which it is based. Therefore, this paper is considered to be preliminary.

Physiographic Setting

Forillon National Park consists of an area of about 240 km², located at the eastern tip of the Gaspé Peninsula, Quebec. It is bounded on the south by Gaspé Bay, on the north by the St. Lawrence River and on the west by Highway 197 and Renard River. Topography is generally mountainous with elevations ranging from sealevel to 550 m.

The area is dominated by a southerly trending part of the Schishock range of the Appalachian Mountain system. The mountains comprise two distinct groups of rocks: one is a sequence of sandstone and limestone of Devonian age that dips uniformly to the southwest and forms a resistant spine or hogback adjacent to Gaspé Bay, and the other is a series of highly folded less resistant schists and phyllites of Ordovician-Cambrian age that occupies the northeastern part of the park and unconformably underlies the Devonian rocks. The northern boundary of the Appalachian Mountains occurs just to the northeast of the park along the St. Lawrence River marked by Logan's Line.

In general, the park is well drained due to its high topographic relief and the occurrence of near-surface bedrock and a thin mantle of dominantly granular soil. Most streams including two larger rivers, the Griffon Cove and Renard Rivers, drain to the St. Lawrence River, whereas a series of smaller but higher gradient streams drain to Gaspé Bay.

Method of Study

Geomorphic features of Forillon Park have been mapped using aerial photographs at a scale of 1:50 000 taken in 1976. This information was transferred directly onto a planimetric base of the area with the aid of a topographic overlay (contour

interval = 50 ft.). In addition, low level oblique colour transparencies of the sea cliffs of the peninsula were examined and all geomorphic features were identified on the film and then located and shown in Figure 1.

All relevant data from a review of principal scientific publications were also transferred onto Figure 1, Especially that compiled by Cumming (1976). Colour slides on file with Parks Canada were examined. These depict principally post-glacial geomorphic features.

Previous Work

Geological work in the park has been reviewed by Cumming (1976). Prest (*in* Douglas 1970), McGerrigle (1952) and Maldague (1973) are the main references for glacial geology. The majority of work in Forillon Park has concentrated on the stratigraphy and paleontology of the Paleozoic bedrock whereas surprisingly little specific data are available on the surficial geology, considerably less on the geomorphology, and nothing on engineering geology.

GEOMORPHIC PROCESSES

The geomorphic processes that have been at work in Forillon National Park can be divided into those that occurred in pre-glacial time, those that occurred in glacial time, and those that are post-glacial or occurring now and immediately following glacial times. Each of these time divisions is discussed below, and the features in the park that represent examples of the processes in each of these time intervals are shown in Figure 1.

Post-glacial

Features formed by post-glacial processes include present coastal erosion and sedimentation, alluvial erosion and sedimentation, colluvium formation, and development of organic terrain. Features formed in immediately post-glacial time (an alpine glacial phase) are included in glacial processes. The possibility of permafrost pockets exists in upland localities, especially in organic depressions as inferred from McGerrigle (1952, p. 48) *in* Cummings (1976, p. 254).

Coastal Processes

The best examples of landforms formed by coastal processes include the formation of a tombola-spit landform at Baie de Penouille and the sea cliffs of Forillon Peninsula.

Development of Baie de Penouille and the associated peninsula is an example of tombola formation by the interaction of coastal processes and alluvial processes. The interpretation presented here is that the tombola formed by deposition of alluvium into the sea (a delta) at the mouth of Ruisseau Ascah which must have interfered with westward longshore drift along Gaspé Bay and caused a

spit to be formed. Eventually the spit grew and merged with the deltaic deposits to form a tombola. The abrupt straight edge on the western part of the tombola may be explained by the unique interaction of tidal currents and river currents from both Dartmouth and York Rivers. This interpretation means that the lagoon is slowly filling due to deltaic sedimentation and overwash along the spit and, in time, the tombola will merge with the shore completely to form a broad sand bar.

Sea cliffs are the most striking geomorphic coastal feature in the area. The coast between Cap Gaspé and Cap Rosier with cliffs up to 200m high is a high wave-energy zone with a wind fetch of 300 km (Owen and Harper 1972). Three basic erosional processes occur along the coast and include undercutting by wave erosion, subsequent rock fall and landslide due to removal of support, and rock fall as a result of frost action. Direct abrasion of the shoreline by shifting and shorefast ice is one aspect of ice erosion, and the other is that material falls onto the ice and is removed by the ice during break-up. The shorefast ice also acts as a defense against wave action.

Wave erosion results in the formation of sea caves and pot holes, differential erosional features, shoreline retreat, and the sedimentation of beach sand and gravel. The beach deposits along Forillon Peninsula are almost entirely due to reworking of talus deposits and, to a very minor degree, to headland erosion of the small resistant outcrops such as Cap Bon Ami.

Along the eastern coast, shoreline erosion is not to be underestimated and should receive considerably more quantitative study. Specifically, it should be possible to determine the rate of erosion using instruments for monitoring present erosion, by the examination and study of air photos and maps, by the analysis of offshore morphology (marine geomorphology), and by interviewing long-time residents of the area.

Erosion by waves and frost along Gaspé Bay is less severe than in the vicinity of Cap Bon Ami and Forillon cliffs. Longshore currents are the most important agents of transport along Gaspé Bay, but the shoreline has been little modified since glacial ice disappeared.

Alluvial Processes

Present alluvial erosion and deposition comprise two main types: that in high gradient streams, and that in low gradient streams. High gradient streams are characterized by rocky bouldery channels, waterfalls, and steep cut banks.

Two streams with particularly well developed cut banks are Ruisseau Beaudry and Ruisseau du Cap des Rosiers where it cuts through moraine as shown in Figure 1). Low gradient streams include Rivière au Renard and Rivière de l'Anse au Griffon (both marked by a shallow meandering type of flood plain in their lower parts but not delineated in Figure 1). Waterfalls are common features.

All streams give rise to small deltas as they enter the sea. The best manifestation of such deltas is the Ruisseau Ascah.

The cutbanks of most streams in the alluvial map unit (No. 1 in Figure 1) in the area may be subject to small slumps. Careful study of these localities is required prior to development of any

kind. Specific setbacks from scarp edges may be required for construction of camp sites adjacent to alluvial terrain.

Organic Deposition

Organic terrain is not common in the park but occurs in poorly drained depressions in lowland and upland localities. The principal process of formation is the anaerobic breakdown of vegetation and gradual accumulation of sphagnum moss. Samples of peat at the base of such terrain may yield material that could be dated and thus provide a minimum age of deglaciation of the area.

All organic and alluvial terrain should be avoided for development due to poor drainage, poor foundation conditions and possible erosion.

Colluvial Processes

Rockfalls and landslides due to undercutting are manifested by large and small piles of broken angular rock that occur in colluvial deposits along slopes and at the base of steep rock faces. Talus fans occur where frost action is an on-going process. Colluvium formed by solid creep is common in the park.

Landslides

Minor rockslides are present along the sea cliff of Forillon Peninsula, and can be expected anywhere the ocean is actively undercutting the cliff.

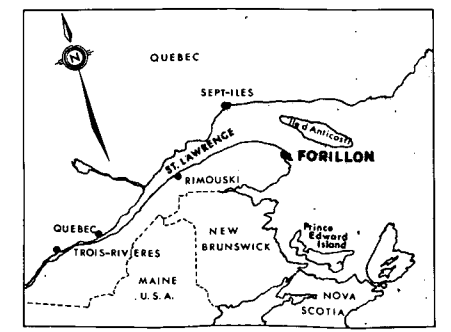
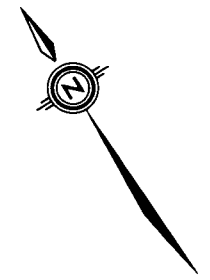
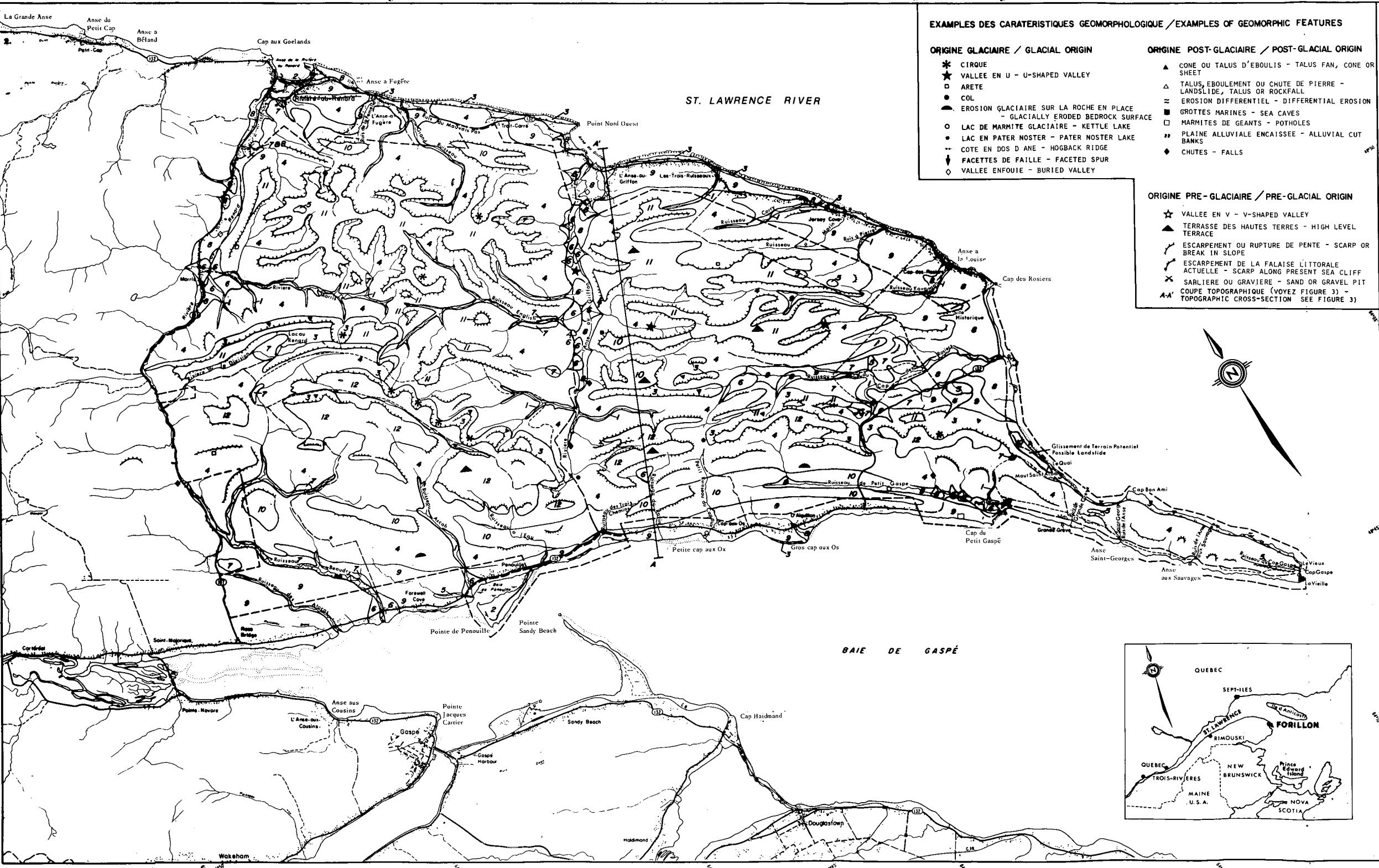
A possible large deep seated landslide may have created the terrace on which the Cap Bon Ami facilities are located. A detailed investigation of this possibility has not been conducted to the writer's knowledge. Certain features of this landform, none of which is conclusive, suggest a landslide origin. First, it is an anomalous landform in an otherwise high steep bedrock shoreline. Its geomorphology is suggestive of a large slump block because its position is compatible with the idea that it may have slipped from a higher position on the face of Mount St. Alban, and thus interpreted as a landslide scar. Russell (*in* Cummings 1976, Vol. 7, p. 273) has mapped faults on both ends of the landform. Other smaller landslides have been identified in the present work. This type of erosion therefore, is characteristic along this stretch of coastline. From an examination of the 9 inch x 9 inch colour negatives (low-level air photos) it is clear that a marked structural discontinuity exists at the base of the slope of this landform. This feature is easily observed in the photos and could be a slip plane of a landslide. Also, it was noted that the strata exposed along the slope of this landform are remarkably similar to strata exposed further to the south along the shoreline - a repetition of strata is therefore suspected and further suggests a landslide origin for this terrace-like feature.

On the other hand, this terrace-like landform is similar to other high level terraces that have been found in the present work, and it could simply be a remnant of one of these features. Also, a number of geologists in the past have examined this feature and none suggested a landslide origin to the writer's knowledge. Most of these geologists, however, were not specifically trained in engineering

PARC NATIONAL FORILLON

Figure 1: Carte Geomorphologique Geomorphie Map

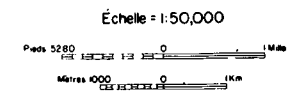
- EXEMPLES DES CARACTERISTIQUES GEOMORPHOLOGIQUE / EXAMLES OF GEOMORPHIC FEATURES**
- ORIGINE GLACIAIRE / GLACIAL ORIGIN**
- * CIRQUE
 - ★ VALLEE EN U - U-SHAPED VALLEY
 - ◻ ARETE
 - ◊ COL
 - EROSION GLACIAIRE SUR LA ROCHE EN PLACE - GLACIALLY ERODED BEDROCK SURFACE
 - LAC DE MARMITE GLACIAIRE - KETTLE LAKE
 - LAC EN PATER NOSTER - PATER NOSTER LAKE
 - ◄ COTE EN DOS D'ANE - HOGBACK RIDGE
 - ⚡ FACETTES DE FAILLE - FACETED SPUR
 - ◇ VALLEE ENFOUIE - BURIED VALLEY
- ORIGINE POST-GLACIAIRE / POST-GLACIAL ORIGIN**
- ▲ CONE OU TALUS D'ÉBOULIS - TALUS FAN, CONE OR SHEET
 - △ TALUS, ÉBOULEMENT OU CHUTE DE PIERRE - LANDSLIDE, TALUS OR ROCKFALL
 - ≡ ÉROSION DIFFÉRENTIEL - DIFFERENTIAL EROSION
 - GROTTES MARINES - SEA CAVES
 - MARMITES DE GEANTS - POTHOLES
 - ⦿ PLAINE ALLUVIALE ENCAISSÉE - ALLUVIAL CUT BANKS
 - ◆ CHUTES - FALLS
- ORIGINE PRE-GLACIAIRE / PRE-GLACIAL ORIGIN**
- ☆ VALLEE EN V - V-SHAPED VALLEY
 - ▲ TERRASSE DES HAUTES TERRES - HIGH LEVEL TERRACE
 - ↘ ESCARPÉMENT OU RUPTURE DE PENTE - SCARP OR BREAK IN SLOPE
 - ↘ ESCARPÉMENT DE LA FALAISE LITTORALE ACTUELLE - SCARP ALONG PRESENT SEA CLIFF
 - × SABLIERE OU GRAVIERE - SAND OR GRAVEL PIT
 - AA' COUPE TOPOGRAPHIQUE (VOYEZ FIGURE 3) - TOPOGRAPHIC CROSS-SECTION (SEE FIGURE 3)



		AGE DES FORMES TERRAIN AGE OF LANDFORM	UNITE DE CARTOGRAPHIE MAP UNIT	FORME DE TERRAIN LANDFORM
QUATÉNAIRE QUATERNARY	RECENT OU HOLOCÈNE RECENT OR HOLOCENE	POST-GLACIAIRE POST-GLACIAL	1	ALLUVIONS ALLUVIAL
			2	PLAGES BEACH
			3	TALUS TALUS
			4	COLLUVIONS COLLUVIAL
			5	SEDIMENTS ORGANIQUES ORGANIC
PLEISTOCÈNE PLEISTOCENE	GLACIAIRE GLACIAL	6	CONE ALLUVIAL SOULÈVÉ RAISED ALLUVIAL FAN	
		7	DEPOT MORAINIQUE MORAINE DEPOSIT	
		8	GLACIO-MARIN GLACIO-MARINE	
TERTIAIRE TO CRÉTACÉ TERTIARY TO CRETACEOUS	PRE-GLACIAIRE PRE-GLACIAL	9	PLACAGE DES BASSES TERRES VENEERED LOWLAND	
		10	BASSES TERRES ONDULÉES ROLLING LOWLAND	
		11	PLACAGE DES HAUTES TERRES VENEERED UPLAND	
		12	HAUTES TERRES DÉBOITÉES BARE UPLAND	

1. NOTE: CARTE PROVISOIRE PRODUITE PAR PHOTO-INTERPRÉTATION, DEMANDE VÉRIFICATION SUR LE TERRAIN.
1. NOTE: PRELIMINARY PHOTO-GEOLOGY INTERPRETATIVE MAP TO BE FIELD CHECKED

2. NOTE: RÉFÉREZ AU TABLEAU 1 POUR L'INFORMATION ADDITIONNELLE
2. NOTE: FOR FURTHER INFORMATION REFER TO TABLE 1



Summary of Geomorphic Units in Forillon National Park
(See Figure 1 for location and distribution of Map Units)

Age of Basic Landform	Map Unit No.	Landform	Composition	Topography	Origin and Past Processes	Present Geomorphic Processes	Engineering Geology
POST-GLACIAL (0 to 12,700 years B.P.)	1	Alluvium	Sand, gravel boulders	High and low gradient gullies and flood plains	Deposited in present stream channels	Fluvial erosion, flood, minor slump	High water table, aggregate, flood and bank stability hazard
	2	Beach	Silt, sand, gravel, lagoonal mud, includes man-made features.	Flat to very gently sloped	Present day shore erosion, deposition includes tombola formation	Active wave erosion of sea cliffs, long-shore tidal and deltaic currents, storm waves, eolain	Active erosion and sedimentation, hazard land, shallow fresh groundwater, possible salty groundwater
	3	Talus	Coarse angular gravel, large blocks in places, bare bedrock	Steep slopes, angle of repose about 30 degrees, bedrock scarps	Active erosion along very steep bedrock scarps, alpine and sea cliff localities	Frost action, wave erosion and undercutting, landslides	Aggregate, rip-rap source, hazard land (rock fall and landslides)
	4	Colluvium	Silty gravel, bare bedrock in places, includes local patches of till and outwash and thin mantle of reworked till	Gently to steeply sloping, some scarps	Soil creep along present slopes and other erosion since area was initially uplifted above sea level.	Present day movement of material down slope due to soil creep processes	Shallow foundations will be unstable, aggregate source, steep slopes
	5	Organic	Peat, organic muck	Closed shallow depressions in uplands and lowlands	Accumulations in poorly drained depressions	Groundwater discharge, springs, slow drainage	Low compressibility, high water table, hazard land
GLACIAL Alpine - 8,000 to 12,000 yrs. B.P. Valley and Continental approximately 10,000 to 12,700 years Before Present	6	Raised Alluvial Fan	Coarse angular silty gravel, may include mud flows, till-like in places	Gentle slope, fanshaped, dissected by streams	Mud flows and melt-water discharge from snowmelt, cirque glaciers and later streams	Inactive today, dissected by present streams, springs	Aggregate source, excellent building site where well-drained, groundwater source
	7	Moraine	Till, sand and gravel, scattered erratics	Rolling to hummocky or hilly	Deposition from glacier	Presently affected by frost action, soil creep, stream erosion	Aggregate source, good foundation, well drained, possible groundwater
	8	Glacio-marine	Marine clay, till and/or outwash overlain by colluvial gravel, includes marine beach gravel in terraces by sea	Flat to gently sloping, minor terraces	Deposited by post-glacial sea in lowland and adjacent valleys adjacent to coast, up to about 75 ft. ASL covered with colluvium and fans in places	Slow deposition of soil creep material from adjacent slopes occurring now, otherwise inactive	Compressible clays, localized poor drainage, best relative agriculture land, beach gravel and colluvium are aggregate sources, groundwater buried valley aquifer

TABLE 1 (Continued)

Age of Basic Landform	Map Unit No.	Landform	Composition	Topography	Origin and Past Processes	Present Geomorphic Processes	Engineering Geology
PRE-GLACIAL Cretaceous to Tertiary erosional Surfaces Affected by all younger processes (i.e. Glaciation)	9	Veneered Lowland	Till and colluvium reworked; bedrock and gravel patches, thin surficial material	Rolling to gently sloping	Pre-glacial erosion of lowland, glaciated by Gaspesian Ice Cap and valley glaciers	Present frost action, erosion along sea cliffs soil creep along slopes, reworked	Aggregate source, poor farmland, good foundation, bedrock near surface
	10	Rolling Lowland	Thin layer of till and residual material over bedrock, reworked, like gravel patches	Flat to gently rolling, bordered by rolling slopes, terrace-like	Pre-glacial erosion, eroded by valley glacier and Gaspesian Ice Cap.	Soil creep along slopes	Soil creep along slopes bedrock near surface
	11	Veneered Upland	Thin layer of colluvium and moraine over bedrock	Gently rolling terrace-like, local depressions	Pre-glacial erosion development of terrace then uplift and dissection by streams, affected by alpine glaciers Gaspesian Ice Cap	Frost action and soil creep presently	Poorly drained in places, some steep slopes, bedrock near surface, deep groundwater aquifers, possible permafrost pockets
	12	Bare Upland	Scattered boulders and granular residual soil, but mainly bare bedrock	Steep slopes, cliffs, rolling depressions	Pre-glacial erosion with development of terrace or peneplain, subsequent uplift and dissection by streams in Tertiary Cretaceous, affected by alpine glaciers and possibly Gaspesian Ice Cap	Frost action and soil creep presently	Bedrock at surface, well drained, steep terrain, scarps, deep groundwater aquifers, possible permafrost pockets

geology and were primarily interested in the stratigraphy and paleontology of the exposed strata. Cumming (1976 p. 189) refers to a gravity slump structure at the subject locality but relates its origin to a Paleozoic disturbance shortly after deposition of the strata; this should not be confused with a possible recent landslide origin for the entire terrace-like landform suggested here.

Talus

Talus is a form of colluvium but since most talus piles or wedge-shaped fans are so distinct and bare of vegetation, they have been identified separately in Figure 1. All terrain mapped as talus (map unit 3 in Figure 1) is subject to rock-fall and is therefore hazardous.

Soil Creep Processes

Colluvium due to soil creep is the most common surficial deposit in the park, and is being formed on most slopes of the area. The process is manifested in road cuts by a layer of residual soil which in some cases resembles till and in many places is stratified. Layers of this soil more or less parallel the slope.

Terrain mapped as colluvium (map unit 4 in Figure 1) is subject to soil creep, in which the upper metre or two of residual soil is slowly moving downhill. Foundations in this material will be unstable, and roads or trails constructed in this unit may be subject to high maintenance costs to remove fallen material and repair warped road surfaces.

Glacial

Glacial features in the park are numerous but few are spectacular. They consist of erosional and depositional forms. Erosional glacial features include a wide variety of alpine landforms such as cirques, cols, faceted spurs, arêtes or knife edge ridges, pater noster lakes, U-shaped valleys and gouged or grooved bedrock. Depositional features include moraine composed of till and (or) outwash sand and gravel, a possible kettle lake in morainal material, raised alluvial fans that were likely deposited by meltwater from small cirques, glaciers, marine clay in the lowlands, and marine gravel in the terraces adjacent to the coast. These features and specific examples of most are shown in Figure 1.

EROSIONAL ORIGIN

Erosional glacial features include cirques, faceted spurs, arêtes, cols, pater noster lakes, grooved or gouged bedrock and U-shaped valleys.

Two levels of cirques are believed to occur: one at an elevation of 274 m (900 feet) and the other at 427 m (1400 feet). The highest cirques are the youngest and most distinct. Lowland cirques and pre-glacial terraces similar to those discussed in this report are known to occur on the west coast of British Columbia adjacent to Hecate Strait.

Faceted spurs, arêtes and cols are numerous in the park and are evidence of erosion by valley glaciers. Pater noster lakes also occur at the head

waters of Ruisseau Ascah. Gouged bedrock such as the steep linear south-facing slope along Gaspé Bay could have only been formed by a massive outlet glacier of the Gaspesian ice cap which likely coalesced with another outlet glacier that flowed down York River.

Large valleys such as that of Griffon Cove River were abraded, overdeepened in part, and carved into U-shaped valleys by glacial ice. These valleys, however, are not classic U-shaped valleys, in fact many of them are more V-shaped and indicative of light glacial erosion compared with, for example, those in the Rocky Mountains and Coast Ranges.

DEPOSITIONAL ORIGIN

Glacial features of depositional origin include low marine terraces, lateral and ground moraine, kame and kettle, and glacial erratics.

Marine clay and marine gravel terraces were formed during a brief period of ocean flooding that occurred immediately following the disappearance of the Gaspesian ice cap. In high energy environments, as in the vicinity of Cap des Rosiers, pebble to cobble beach gravel was deposited in terraces; in sheltered but flooded valleys, such as Griffon Cove River valley, marine clay was deposited. Numerous marine shells can be found in this clay in places and these have been dated. Near the mouth of Griffon Cove River, the Carbon 14 method was employed to date shells found there. An age of 12,400 years B.P. was obtained (Prest, personal communication, 1978). This means that no glacial ice existed and also, that the St. Lawrence River was likely free of ice at that time. This age thus provides a date for the end of the continental phase of Wisconsin glaciation in this region. Dr. Prest also found shells (dated at 12,700 years B.P.) in marine clay 1.6 km south of Rivière au Renard near the park boundary. The clay occurs as an eroded terrace overlain in places by outwash gravel which indicates an ice advance after marine invasion at this locality.

Till and outwash gravel occur in ground and (or) lateral moraine of the park. The till is very silty and clayey and has a wide variety of sizes of pebbles and boulders. Gravel deposits are not numerous.

A possible kettle lake occurs just below and south of the crest of Mount St. Alban. Another one occurs just to the west of the park boundary at an elevation of 274 m (900 feet) 1.6 km east of Morris.

Although glacial erratics of granite composition from a Canadian Shield source to the north have not been found in the park to date, which means that the Laurentide continental glacier did not reach the park, the most favourable place to search for such erratics would be in the uplands of the northeastern part of the park. Here a very thin layer of till is thought to occur in places overlying the bedrock. In many localities the till may be represented only by scattered stones.

Pre-glacial Features

Pre-glacial geomorphic features are well represented in the park but are somewhat obscure to a casual observer because of the scale of the features. These include a series of erosional terraces that

SUMMARY AND CONCLUSIONS

compose the uplands, midlands and lowlands of the park. The terraces have all been modified by ice cap and alpine glaciation, and colluvial processes, but basically consist of a gently rolling surface flanked by a scarp or break in slope that gives way to a steep slope. The surfaces of the terraces are poorly defined in some places but four surfaces, as listed below, have been recognized in the present work and all appear to have been tilted or warped in some manner.

Pre-glacial Terraces	Elevation of Surface (in feet)
Map Unit 12	1000 rising NW to 1800
Map Unit 11	700 rising NW to 1250
Map Unit 10	300 rising NW to 1000
Map Unit 9	100 rising NW to 500

The terraces may represent successive levels of erosion that were attained in certain erosional intervals during Cretaceous and Tertiary time. That is, for example, the highest level was at one time sealevel in the area, and has since been uplifted, and so on, down to the lowest terrace level. The combined perspective and topographic cross-section of Figure 2 is provided to help illustrate these features.

Very little is known about these terraces, however, the opinion expressed here is that these terraces are ancestral to the uplifting and tilting of the continental margin in the Cretaceous-Tertiary portion of pre-glacial time (see also Mathews 1974).

The other visible feature related to pre-glacial geomorphic development is the present stream pattern which is thought to be largely pre-glacial in origin, and only slightly modified by glacial processes. The main manifestation of pre-glacial origin is the way the major valleys of Griffon Cove River and Renard River have been truncated by glacial erosion along Gaspé Bay, and the neat V-shape of some of the other valleys as opposed to a distinct U-shaped glacially eroded valley.

The broad geologic events that have affected Forillon Park are summarized below. This provides a framework for identification of certain geomorphic features shown in Figure 1 that relate to various past and present events in the park. A summary of geomorphic features is given in Table 1 which should be used in conjunction with Figure 1.

The geomorphic evolution of Forillon National Park begins in late Precambrian time when the ancient Laurentian Mountains dominated this part of the land. Then the Acadian continental marginal basin began to develop in early Cambrian time. Land-derived detritus was poured into this basin throughout Cambrian and Ordovician times and resulted in a very thick accumulation of mainly limestone - shale and carbonate sands (St. Julien and Hubert 1975). The instability of the earth's crust then increased considerably, which resulted in intense folding, faulting, and metamorphism. This was the Taconic Orogeny and was so intense that large slabs of allocthonous schistose rock, or klippen, were physically dislocated and shoved many kilometres to the north (St. Julien and Hubert 1975). Also, gravity sliding occurred but nappes are not present in the park. The schistose rocks that outcrop in the northeastern part of the park belong to klippen of the allocthonous sequence.

A long period of erosion occurred throughout much of Silurian time. Then in Devonian time another sedimentation basin developed along the modern Appalachian Mountains (Williams *in* Price and Douglas 1972). Thick sequences of limestone and shale were deposited and volcanic activity occurred in places. Towards the end of Devonian time this basin was deformed during the Acadian Orogeny. Also, the rocks were intruded by granite at Mt. McGorrigle and, finally, the region became a stable rising landmass.

The Appalachian landmass was at one time mountainous, but erosion reduced the region to nearly a flat plain. For example, all of the Paleozoic rocks were stripped from the Laurentian Mountains area. This erosion occupied the time between the late Devonian and the Jurassic Periods, some 300 million years, and the erosional plain that formed

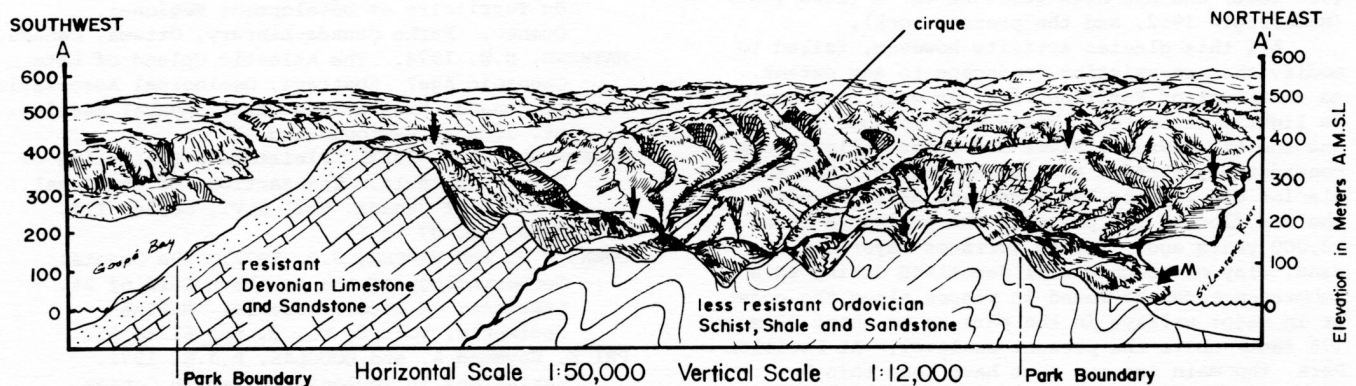


FIG. 2 Four gently rolling erosional surfaces (poorly defined terraces) can be distinguished in Forillon Park as depicted in this sketch or perspective along topographic cross-section A-A' (location shown in Fig. 1). Arrows indicate the various levels - these are believed to be related to periods of uplift, erosion and gradual dissection of the Atlantic Penepain. A low terrace designated M related to post-glacial marine invasion is situated 15 to 23 meters above sea level and should not be confused with the older high terraces. Two levels of cirques are also depicted in this artist's sketch.

has been called the Atlantic Peneplain (King *in* Price and Douglas 1972, p. 252).

In Jurassic time the Atlantic Ocean as we know it began to form between North America and Europe. This caused uplift of the Atlantic Peneplain and tilting of it toward the south. Therefore, throughout much of the Cretaceous and Tertiary Periods the Gaspé Peninsula area and the Canadian shield to the north were subject to further erosion whereas in the offshore part of the coast (the modern Gulf of St. Lawrence) subsidence occurred and sediment derived from the Gaspé and other land areas was deposited. The tops of the highest hills in Forillon Park, as well as, although modified by fluvial erosion during successive intervals, by glacial erosion, may represent ancestral remnants of the Atlantic Peneplain.

Break-up and drifting of the continents, which is still going on today, continued and was accompanied first in the Early Cretaceous by igneous intrusions, some of which occur in the Monteregian Hills, and then by faulting. A fault graben may have developed along the St. Lawrence at this time, according to some authors, and this could have created the ancestral valley of the modern St. Lawrence River (Kumarapelli 1976).

By Quaternary time the basic land features consisting of terraced uplands and lowlands, and V-shaped valleys in Forillon National Park had been created. Although four major glacier advances took place during the Pleistocene Epoch of the Quaternary Period, at Forillon there is evidence for only one ice cap that probably engulfed the area. This was the Gaspesian Ice Cap and it spread from central Gaspé about 100,000 years ago during Wisconsinan time. Gaspesian ice presumably met the Laurentide continental glacier along the St. Lawrence River Valley, perhaps very close to the northern boundary of Forillon Park. This glacial event is considered to be classical Wisconsin in age. The Gaspesian Ice Cap phase was likely preceded by cirque and valley glacier development.

After the Gaspean ice cap had melted, valley glaciers developed in part of the Gaspé Peninsula. In Forillon there is evidence for at least two cirque glacier phases, one low level at 274m (900 feet) and one high level at 427 m (1400 feet) (McGerrigle 1952, and the present work).

All this glacier activity however, failed to modify the pre-existing landscape to any extent, so that most of Forillon Park can only be described as lightly glaciated. Hills were rounded somewhat and the St. Lawrence River and other valleys were considerably abraded and over-deepened. But the glacier did depress the land, and as it receded the sea flooded low parts of the peninsula about 12,000 years ago. Marine terraces layered with sand, clay or beach gravel deposited during this submergence can be found in places along the coast or in major valleys in the park up to about 23 m (75 feet) above the present sealevel. At Forillon Park, the main ice cap must have been thin because post-glacial ocean flooding occurred only up to about 23 m (75 feet) above sea level (Prest *in* Douglas 1970), whereas further south evidence for this marine invasion up to several hundred metres exists. Thus the Gaspesian Ice Cap was either much thicker to the south, or the land adjacent to the sea at Forillon Park was still

occupied by ice when the marine invasion took place. The present shoreline has therefore been subjected to the full force of the sea throughout the last 8,000 to 9,000 years probably, and the spectacular cliffs of Forillon Peninsula were formed during this time.

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